

Tutorial on data assimilation

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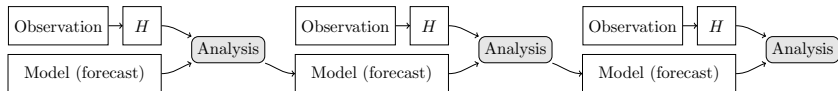


Outline

- 1 Data assimilation: principles
 - Definition
 - Mathematical framework
- 2 Main techniques
 - 3D-Var and optimal interpolation
 - The Kalman filter
 - The ensemble Kalman filter
 - 4D-Var
- 3 Advanced techniques
 - Ensemble variational methods
 - Particle filters
- 4 Uncertainty quantification of the best estimate

Data assimilation: definition

- **Data assimilation** is the set of techniques that allow to **optimally** combine **observations** of a physical system with **numerical models** and **prior information** of that system, so as to get an estimate of the state of the system.
- In the geosciences: Numerical models are often computationally costly. They are often dynamical.
- In the geosciences: The state space and observations space are huge (up to $10^9/10^7$ for operational systems, up to $10^7/10^5$ for research systems). A big data problem with costly dynamical models.
- What for?: estimate initial state of chaotic systems for forecasting, re-analysis, estimate parameters (\sim inverse modelling).
- Example: Data assimilation for prediction.



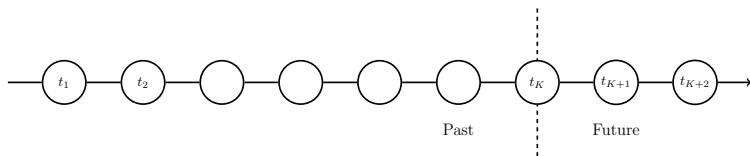
Data assimilation system

- Data assimilation system = observation and evolution models + statistics of the errors. Typically:

$$\mathbf{x}_k = M_{k:k-1}(\mathbf{x}_{k-1}) + \eta_k$$

$$\mathbf{y}_k = H_k(\mathbf{x}_k) + \varepsilon_k$$

with $\eta_k \sim \mathcal{N}(\mathbf{0}, \mathbf{Q}_k)$ and $\varepsilon_k \sim \mathcal{N}(\mathbf{0}, \mathbf{R}_k)$.



- Denoting $\mathbf{x}_{K:1} = \mathbf{x}_1, \mathbf{x}_2, \dots, \mathbf{x}_K$, $\mathbf{y}_{K:1} = \mathbf{y}_1, \mathbf{y}_2, \dots, \mathbf{y}_K$:
- Prediction: Estimate \mathbf{x}_k for $k > K$, knowing $\mathbf{y}_{K:1}$;
 - Filtering: Estimate \mathbf{x}_K , knowing $\mathbf{y}_{K:1}$;
 - Smoothing: Estimate $\mathbf{x}_{K:1}$, knowing $\mathbf{y}_{K:1}$.

The ideal mathematical framework

- ▶ Bayes/Laplace approach:

$$p(\mathbf{x}|\mathbf{y}) = \frac{p(\mathbf{y}|\mathbf{x})p(\mathbf{x})}{p(\mathbf{y})}$$

with $p(\mathbf{y}|\mathbf{x})$ the **likelihood** of the observations, $p(\mathbf{x})$ the **prior**/background on the system's state, and $p(\mathbf{y})$ the **evidence**. The evidence is a normalisation that does not depend on \mathbf{x} :

$$p(\mathbf{y}) = \int d\mathbf{x} p(\mathbf{y}|\mathbf{x})p(\mathbf{x}).$$

- ▶ This is a probabilistic approach. It quantifies the uncertainty/the information. It does not provide a deterministic **estimator**. This would require to make a choice on top of Bayes' rule.
- ▶ The Bayesian approach is very satisfactorily [Jaynes et al., 2003]. Most DA methods can be derived or comply with Bayes' rule.
- ▶ But it does not lend to a closed form analytically tractable solution.

Gaussian approximation

- ▶ A key to obtain a (approximate) solution is to truncate the errors to second-order moments \sim [the Gaussian approximation](#). Most of DA methods are fully or partially based on this assumption.
- ▶ The elementary building block of DA schemes is the statistical [BLUE](#) (for Best Linear Unbiased Estimator) analysis. Time is considered fixed. \mathbf{H} is assumed linear.

$$\mathbf{y} = \mathbf{H}\mathbf{x} + \varepsilon^o \quad \mathbf{x} = \mathbf{x}^b + \varepsilon^b$$

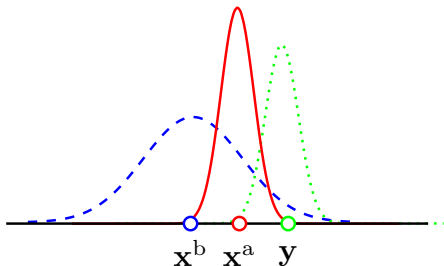
where $\varepsilon^o \sim \mathcal{N}(\mathbf{0}, \mathbf{R})$, and $\varepsilon^b \sim \mathcal{N}(\mathbf{0}, \mathbf{B})$.

- ▶ Solution:

$$\mathbf{x}^a = \mathbf{x}^b + \mathbf{K}(\mathbf{y} - \mathbf{H}\mathbf{x}^b)$$

$$\mathbf{K} = \mathbf{B}\mathbf{H}^T(\mathbf{R} + \mathbf{H}\mathbf{B}\mathbf{H}^T)^{-1}$$

$$\mathbf{P}^a = (\mathbf{I} - \mathbf{K}\mathbf{H})\mathbf{B}.$$



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3D-Var and optimal interpolation

- Variational formulation of the same problem

$$J(\mathbf{x}) = \frac{1}{2} \|\mathbf{x} - \mathbf{x}^b\|_{\mathbf{B}^{-1}}^2 + \frac{1}{2} \|\mathbf{y} - \mathbf{H}\mathbf{x}\|_{\mathbf{R}^{-1}}^2$$

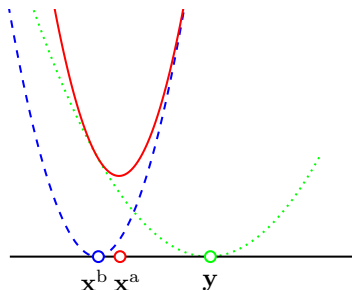
where $\|\mathbf{x}\|_{\mathbf{A}}^2 = \mathbf{x}^T \mathbf{A} \mathbf{x}$, which is equivalent to BLUE.

- Probabilistic/Bayesian interpretation:

$$p(\mathbf{x}|\mathbf{y}) \propto e^{-J(\mathbf{x})}$$

- Capable of handling nonlinear observation operator using standard nonlinear optimisation methods:

$$J(\mathbf{x}) = \frac{1}{2} \|\mathbf{x} - \mathbf{x}^b\|_{\mathbf{B}^{-1}}^2 + \frac{1}{2} \|\mathbf{y} - H(\mathbf{x})\|_{\mathbf{R}^{-1}}^2.$$



Chaining the analyses in time

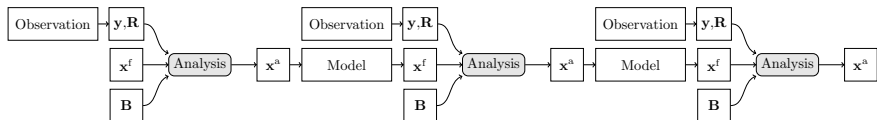
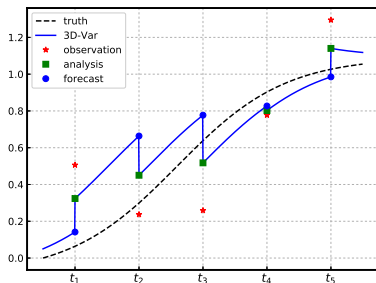
► Chaining the BLUE/3D-Var cycles:

- ① Analysis with a forecast at t_k : \mathbf{x}_k^f and with static information \mathbf{B} : \mathbf{x}_k^a ;
- ② Forecast to t_{k+1} : $\mathbf{x}_{k+1}^f = M_{k+1:k}(\mathbf{x}_k^a)$.

► Also known as **optimal interpolation** (if the analysis step is BLUE).

► Relatively cheap. Used in oceanography, atmospheric chemistry. Requires a smart construction of \mathbf{B} .

► But the information about the errors is not propagated in time ...



The Kalman filter

► Similar to optimal interpolation. But, now, we want to replace the static \mathbf{B} with a dynamic \mathbf{P}^f which needs updating and propagating.

► Analysis step:

$$\mathbf{x}_k^a = \mathbf{x}_k^f + \mathbf{K}_k \left(\mathbf{y}_k - \mathbf{H}_k \mathbf{x}_k^f \right),$$

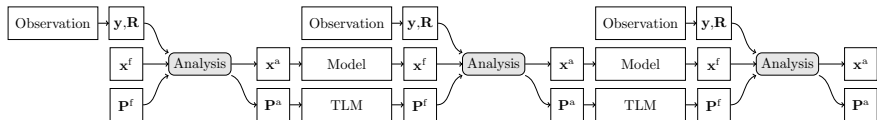
$$\mathbf{K}_k = \mathbf{P}_k^f \mathbf{H}_k^T \left(\mathbf{R}_k + \mathbf{H}_k \mathbf{P}_k^f \mathbf{H}_k^T \right)^{-1},$$

$$\mathbf{P}_k^a = \left(\mathbf{I} - \mathbf{K}_k \mathbf{H}_k \right) \mathbf{P}_k^f.$$

► Forecast step:

$$\mathbf{x}_{k+1}^f = \mathbf{M}_{k+1:k} \mathbf{x}_k^a,$$

$$\mathbf{P}_{k+1}^f = \mathbf{M}_{k+1:k} \mathbf{P}_k^a \mathbf{M}_{k+1:k}^T + \mathbf{Q}_{k+1}.$$



The extended Kalman filter

- ▶ Optimal if the models are linear and if all the initial and observations errors are Gaussian: it gives the perfect Gaussian solution of Bayes' rule.
- ▶ Can be extended to nonlinear models: then

$$\begin{aligned}\mathbf{x}_{k+1}^f &= M_{k+1:k}(\mathbf{x}_k^a), \\ \mathbf{P}_{k+1}^f &= \mathbf{M}'_{k+1:k} \mathbf{P}_k^f \mathbf{M}'_{k+1:k}{}^T + \mathbf{Q}_{k+1},\end{aligned}$$

where $\mathbf{M}'_{k+1:k}$ is the tangent linear model.

- ▶ Extremely costly for large geophysical models: storage space (storage of \mathbf{P}^f) and computations ($\mathbf{M}'_{k+1:k} \mathbf{P}_k^f \mathbf{M}'_{k+1:k}{}^T$ requires $2n$ integrations of the model).

The ensemble Kalman filter

- ▶ The idea [Evensen, 1994; Houtekamer and Mitchell, 1998] is to make the KF work in high dimensions and replace \mathbf{P} (\mathbf{P}^a or \mathbf{P}^f) with an ensemble of states $\mathbf{x}_1, \mathbf{x}_2, \dots, \mathbf{x}_m$. The moments of the error could theoretically be approximated by [the sample/empirical moments](#):

$$\mathbf{x}^f = \frac{1}{m} \sum_{i=1}^m \bar{\mathbf{x}}, \quad \mathbf{P}^f = \frac{1}{m-1} \sum_{i=1}^m (\mathbf{x}^{(i)} - \bar{\mathbf{x}}) (\mathbf{x}^{(i)} - \bar{\mathbf{x}})^T.$$

- ▶ Analysis step: Similar to the KF but \mathbf{P}^f explicitly or implicitly taken as the [sample covariance estimator](#).
- ▶ Forecast step: The ensemble is propagated using the full nonlinear model (not the tangent linear model!)

$$\mathbf{x}_{k+1}^{(i),f} = M_{k+1:k} (\mathbf{x}_k^{(i),a}).$$

The ensemble Kalman filter

- ▶ Two main flavors of EnKFs: **stochastic and deterministic**, but many variants.
- ▶ The stochastic EnKF is the closest to traditional KF, but adds stochastic perturbations to the observations of each members to properly account for the observation errors [Burgers et al., 1998]:

$$\mathbf{x}_{(i)}^a = \mathbf{x}_{(i)}^f + \mathbf{K} \left(\mathbf{y} + \varepsilon_{(i)} - \mathbf{H}\mathbf{x}_{(i)}^f \right).$$

- ▶ The deterministic EnKF avoids the introduction of the stochastic perturbations by updating the square root of $\mathbf{P}^f = \mathbf{X}_f \mathbf{X}_f^T$, i.e. \mathbf{X}_f . One of the variant (ETKF, [Hunt et al., 2007]) operates the linear algebra **in the space of the perturbations**:

$$\mathbf{x}^a = \mathbf{x}^f + \mathbf{X}_f \mathbf{w}^a.$$

The analysis in the perturbation space is given by

$$\mathbf{w}^a = \left(\mathbf{I}_m + \mathbf{Y}_f^T \mathbf{R}^{-1} \mathbf{Y}_f \right)^{-1} \mathbf{Y}_f^T \mathbf{R}^{-1} \left(\mathbf{y} - \mathbf{H}\mathbf{x}^f \right)$$

where $\mathbf{Y}_f = \mathbf{H}\mathbf{x}^f$. This updates the mean state via $\mathbf{x}^a = \mathbf{x}^f + \mathbf{X}_f \mathbf{w}^a$. The perturbations around it are updated via

$$\mathbf{X}_a = \mathbf{X}_f \left(\mathbf{I}_m - \mathbf{Y}_f^T (\mathbf{Y}_f \mathbf{Y}_f^T + \mathbf{R})^{-1} \mathbf{Y}_f \right)^{\frac{1}{2}} \mathbf{U}, \quad \text{where } \mathbf{U} \in O(N) \quad \text{and} \quad \mathbf{U}\mathbf{1} = \mathbf{1}.$$

The downside of the EnKF: rank-deficiency

- There is a heavy price to pay for replacing the \mathbf{P}^f $n \times n$ covariance matrix of the KF with the \mathbf{X}_f $m \times n$ anomaly matrix: **spurious correlations** for distant state components. If $\mathbf{P} = \mathbf{X}_f \mathbf{X}_f^T$ and \mathbf{B} is the true error covariance matrix of a Gaussian process:

$$\text{Cov}([\mathbf{P}]_{ii}, [\mathbf{P}]_{jj}) = \frac{2}{N-1} [\mathbf{B}]_{ij}^2, \quad \text{Cov}([\mathbf{P}]_{ij}, [\mathbf{P}]_{ij}) = \frac{1}{N-1} \left([\mathbf{B}]_{ij}^2 + [\mathbf{B}]_{ii} [\mathbf{B}]_{jj} \right).$$

- But, for geophysical systems, we know that most long-range correlations are dampened exponentially. Consequently, the covariances are misestimated (too low variances, too high long-range covariances) and leads to divergence of the EnKF. —→ Practically, this is solved using two fixes: **inflation** and **localisation**.
- Inflation consists in inflating the covariances by a scalar in the hope to compensate for the underestimation of the error statistics [Pham et al., 1998, Anderson et al., 1999]:

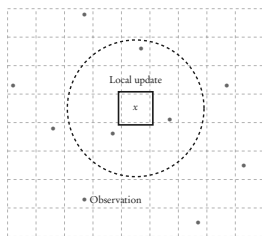
$$\mathbf{x}_{(i)} \longleftarrow \mathbf{x}_{(i)} + \lambda \left(\mathbf{x}_{(i)} - \bar{\mathbf{x}} \right).$$

Can be avoided in a perfect-model context: finite-size EnKF (EnKF-N) [Bocquet et al., 2011-2018].

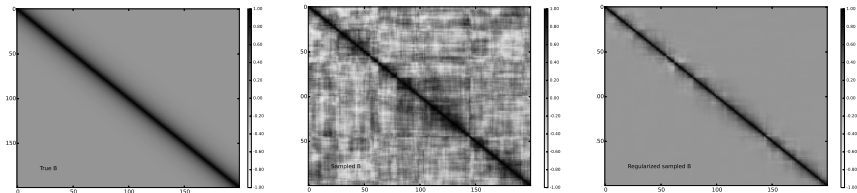
Localisation

- ▶ Two flavors of localisation: **domain localisation** and **covariance localisation**.

- ▶ **Domain localisation**: divide and conquer.
The DA analysis is performed in parallel in local domains. The outcomes of these analyses are later sewed together. This is applicable only if the long-range error correlations are negligible.

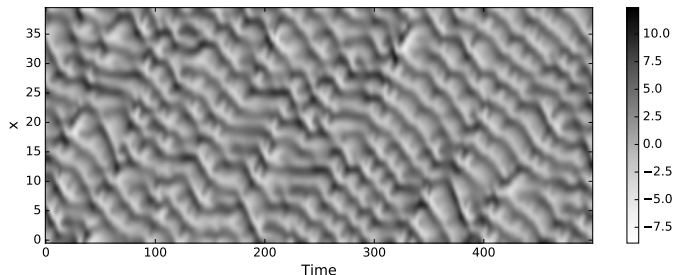


- ▶ **Covariance localisation**: killing off spurious correlation explicitly: $\mathbf{P}^f = \rho \circ (\mathbf{X}_f \mathbf{X}_f^T)$.



- ▶ These strategies have successfully been applied to the EnKF [Hamill et al, 2001; Houtekamer and Mitchell, 2001; Evensen, 2003; Hunt et al., 2007].

Nonlinear chaotic models: the Lorenz-95 low-order model



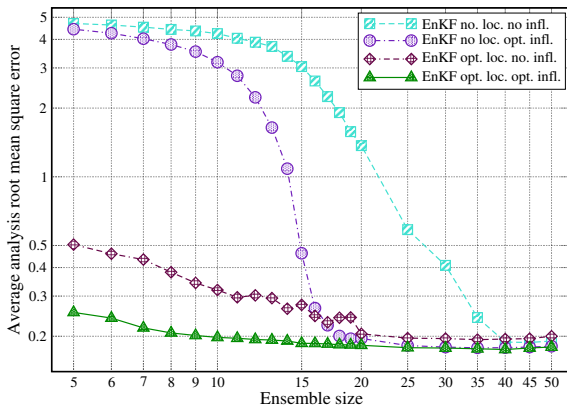
- It represents a mid-latitude zonal circle of the global atmosphere.
- Set of $n = 40$ ordinary differential equations [Lorenz and Emmanuel 1998]:

$$\frac{dx_i}{dt} = (x_{i+1} - x_{i-2})x_{i-1} - x_i + F, \quad (1)$$

where $F = 8$, and the boundary is cyclic.

- Conservative system except for a forcing term F and a dissipation term $-x_i$.
- Chaotic dynamics, 13 positive and 1 neutral Lyapunov exponents, a doubling time of about 0.42 time units.

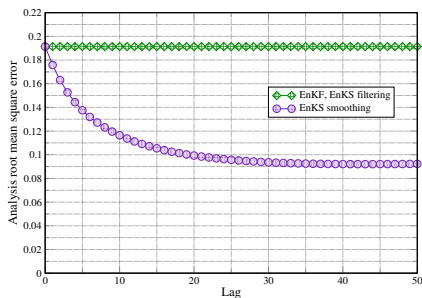
Illustration with the Lorenz-95 model



► Performance of the EnKF in the absence/presence of inflation/localisation.

What about smoothing?

- ▶ There are smoothing variants of the Kalman filter [Anderson & Moore, 1979], the Kalman smoother used in the geosciences [Cohn et al., 1994]
- ▶ And they have been adapted to the EnKF and variants [Evensen & van Leeuwen, 2000], [Evensen, 2009], [Cosme et al., 2012], [Bocquet & Sakov, 2014], etc.
- ▶ Sometimes called asynchronous data assimilation [Sakov et al., 2010; Sakov & Bocquet, 2018].
- ▶ With the notable exception of the IEnKS, these smoothers relies on Gaussian assumptions within the DAW.
- ▶ 4D-Var is a more natural method to handle nonlinearity within the DAW.



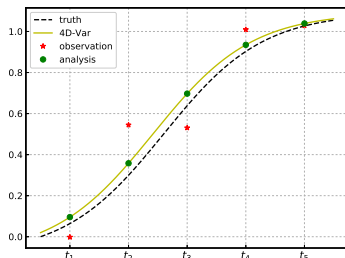
4D-Var

- ▶ Strongly constrained 4D-Var, i.e. assuming the model is perfect

$$J(\mathbf{x}_0) = \frac{1}{2} \|\mathbf{x}_0 - \mathbf{x}_0^{\mathbf{b}}\|_{\mathbf{B}^{-1}}^2 + \frac{1}{2} \sum_{k=1}^K \|\mathbf{y}_k - H_k(\mathbf{x}_k)\|_{\mathbf{R}_k^{-1}}^2$$

under the constraints that $\mathbf{x}_{k+1} = M_{k+1:k}(\mathbf{x}_k)$ for $k = 0, \dots, K-1$.

- ▶ Fits a model trajectory through the 4D data points.
- ▶ In high-dimensional spaces, requires $\nabla_{\mathbf{x}_0} J$ for an efficient minimisation. But $\nabla_{\mathbf{x}_0} J$ depends on the adjoint of $M_{k+1:k}$ and H_k . This can be a very difficult technical task if the model is a huge piece of code for a nonlinear high-dimensional model.



- ▶ Weakly constrained 4D-Var, i.e. assuming the model is imperfect

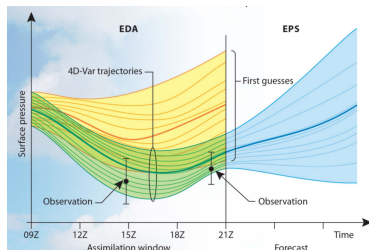
$$J(\mathbf{x}_{K:0}) = \frac{1}{2} \|\mathbf{x}_0 - \mathbf{x}_0^{\mathbf{b}}\|_{\mathbf{B}^{-1}}^2 + \frac{1}{2} \sum_{k=0}^K \|\mathbf{y}_k - H_k(\mathbf{x}_k)\|_{\mathbf{R}_k^{-1}}^2 + \frac{1}{2} \sum_{k=1}^K \|\mathbf{x}_k - M_{k:k-1}(\mathbf{x}_{k-1})\|_{\mathbf{Q}_k^{-1}}^2.$$

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Hybridising ensemble and variational methods

- ▶ Hybrid: Use flow-dependent statistics from an EnKF into 3D-Var [Hamil & Snyder 2000; Wang et al. 2007].
- ▶ 4D-LETKF [Hunt et al., 2004; Fertig et al., 2007]
- ▶ EDA: ensemble of 4D-Var (ECMWF, Météo-France) [Raynaud et al., 2009; Bonavita et al., 2012; Berre et al., 2015; Jardak & Talagrand 2018]



- ▶ 4D-EnVar: Adjoint-less 4D-Var [Liu et al., 2008; Buehner et al. 2010; Zhang and Zhang, 2012; Fairbairn et al. 2014, Desroziers et al. 2014], but ensemble update and nonlinearity still not completely addressed.
- ▶ IEnKS: has it all [Sakov et al. 2012, Bocquet & Sakov 2012-2016].
- ▶ As ensemble methods, they all require localisation, which is more difficult to implement in a 4D context [Bocquet, 2016] except if the adjoint is available.
 - For a review on EnVar methods, see Chapter 7 of the new book [Asch et al., 2016].

The iterative ensemble Kalman smoother (IEnKS)

- ▶ Reduced scheme in ensemble space, $\mathbf{x}_0 = \bar{\mathbf{x}}_0 + \mathbf{X}_0 \mathbf{w}$, where \mathbf{X}_0 is the ensemble anomaly matrix:

$$\tilde{J}(\mathbf{w}) = J(\bar{\mathbf{x}}_0 + \mathbf{X}_0 \mathbf{w}).$$

- ▶ Analysis IEnKS cost function in ensemble space:

$$\tilde{J}(\mathbf{w}) = \frac{1}{2} \sum_{k=1}^L \|\mathbf{y}_k - H_k \circ M_{k:0}(\bar{\mathbf{x}}_0 + \mathbf{X}_0 \mathbf{w})\|_{\beta_k \mathbf{R}_k^{-1}}^2 + \frac{1}{2} (N-1) \|\mathbf{w}\|^2.$$

$\{\beta_0, \beta_1, \dots, \beta_L\}$ weight the observations impact within the window.

- ▶ As a variational **reduced** method, one can use Gauss-Newton [Sakov et al., 2012], Levenberg-Marquardt [Bocquet & Sakov, 2012], quasi-Newton, trust region, etc., minimisation schemes.

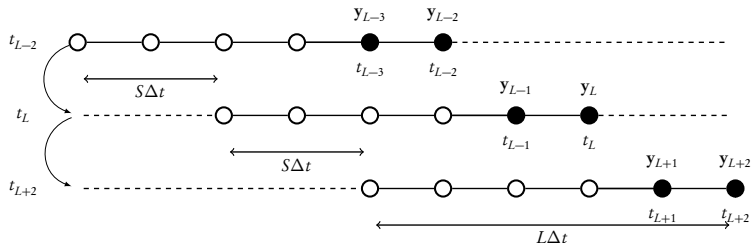
- ▶ **Perturbation update**: same as the ETKF

$$\mathbf{E}_0^* = \mathbf{x}_0^* \mathbf{1}^T + \sqrt{N-1} \mathbf{X}_0 \left[\nabla_{\mathbf{w}}^2 \tilde{J} \right]_*^{-1/2} \mathbf{U} \quad \text{where} \quad \mathbf{U} \in \mathbf{O}(N) \quad \text{and} \quad \mathbf{U} \mathbf{1} = \mathbf{1}.$$

→ Cécile Defforge's talk.

Chaining 4D analyses in time

- ▶ The IEnKS opens up new perspectives on the chaining of DA cycles which was little relevant for either the EnKF or 4D-Var.
- ▶ L : length of the data assimilation window,
- ▶ S : shift of the data assimilation window in between two updates.



Variational analysis in ens. space

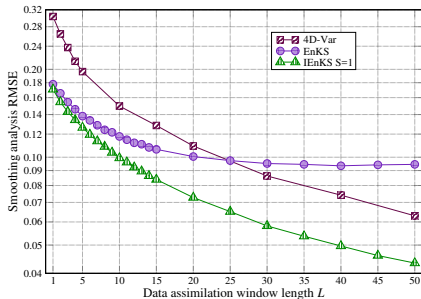
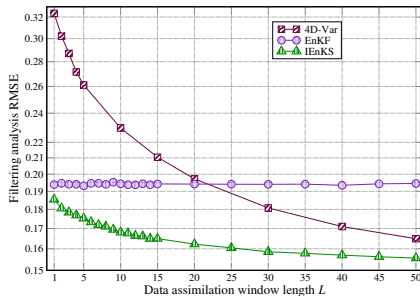
→

Posterior ens. generation

→

Ens. forecast

Performance comparison with Lorenz-95



- Comparing 4D-Var, the EnKF, the EnKS and the IEnKS.

Taking the bull by the horns: the particle filter

- ▶ The particle filter is the **Monte-Carlo solution of the Bayes' equation**. This is a sequential Monte Carlo method.
- ▶ The most simple algorithm of Monte Carlo type that solves the Bayesian filtering equations is called the **bootstrap particle filter** [Gordon et al. 1993] .

Sampling: Particles $\{\mathbf{x}_1, \mathbf{x}_2, \dots, \mathbf{x}_M\}$.
 Pdf at time t_k : $p_k(\mathbf{x}) \simeq \sum_{i=1}^M \omega_i^k \delta(\mathbf{x} - \mathbf{x}_k^i)$.

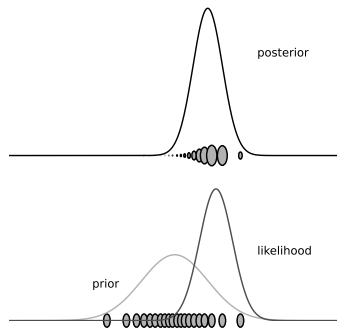
Forecast: Particles propagated by

$$p_{k+1}(\mathbf{x}) \simeq \sum_{i=1}^M \omega_i^k \delta(\mathbf{x} - \mathbf{x}_{k+1}^i)$$

with $\mathbf{x}_{k+1}^i = \mathbf{M}_{k+1}(\mathbf{x}_k^i)$.

Analysis: Weights updated according to

$$\omega_{k+1}^{a,i} \propto \omega_{k+1}^{f,i} p(\mathbf{y}_{k+1} | \mathbf{x}_{k+1}^i).$$



- ▶ Analysis is carried out with only a few multiplications. No matrix inversion!

Taking the bull by the horns: the particle filter

► These normalised statistical weights have a potentially large amplitude of fluctuation. One particle (one trajectory of the model) will stand out among the others. Its weight will largely dominate the others ($\omega_i \lesssim 1$). Then the particle filter becomes very inefficient as an estimating tool since it has lost its variability. This phenomenon is called **degeneracy** of the particle filter [Kong et al. 1994].

Resampling One way to mitigate this phenomenon is to resample the particles by redrawing a sample with uniform weights from the degenerate distribution. After resampling, all particles have the same weight: $\omega_k^i = 1/M$.

► Handles very well, very nonlinear low-dimensional systems. But, without modification, very inefficient for high-dimensional models. Avoiding degeneracy requires a great number of particles that scales exponentially with the size of the system. This is a manifestation of the **curse of dimensionality**.

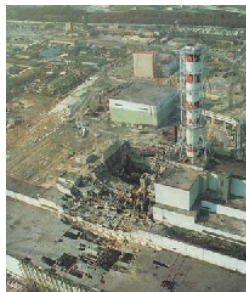
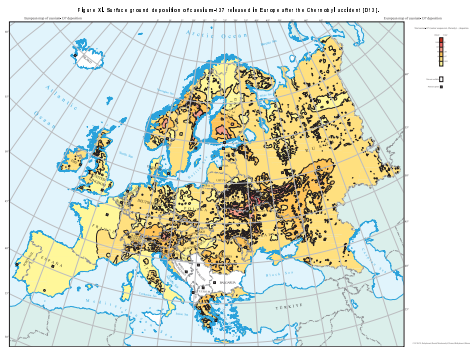
Application of the particle filter in the geosciences

- ▶ The applicability of particle filters to high-dimensional models has been investigated in the geosciences [van Leeuwen, 2009; Bocquet, 2010]. The impact of the curse of dimensionality has been quantitatively studied in [Snyder et al., 2008]. It was known [Mackay et al., 2003] that using an importance proposal to guide the particles towards regions of high probability will not change this trend, albeit with a reduced exponential scaling, which was confirmed by [Snyder et al., 2015]: [optimal importance sampling particle filter](#) [Doucet et al., 2000; Bocquet, 2010; Snyder; 2011].
- ▶ Particle smoother over a data assimilation window, alternative and more efficient particle filters can be designed, such as the [implicit particle filter](#) [Morzfeld et al., 2012].
- ▶ Particle filters can nevertheless be useful for high-dimensional models if the significant degrees of nonlinearity are confined to a small subspace of the state space, e.g. Lagrangian data assimilation [Slivinski et al., 2015] .
- ▶ It is possible possible to design nonlinear filters for high-dimensional models such as the [equal-weight particle filter](#) [van Leeuwen & Ades, 2010-2017].
- ▶ Localisation can be (should be?) used in conjunction with the particle filter [Reich et al. 2013; Potterjoy, 2016; Penny & Miyoshi, 2016; Farchi & Bocquet, 2018].
→ Alban Farchi's talk.
- ▶ It has been applied in hydrology, nivology, climate, etc [Goosse, Dubinkina et al.].

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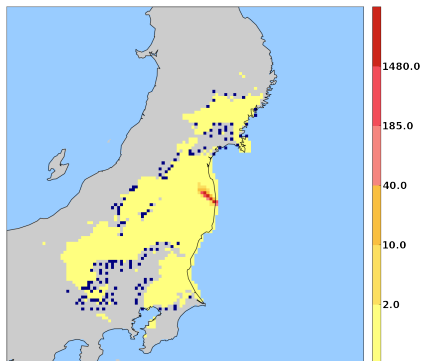
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- 4 Uncertainty quantification of the best estimate

Case study: Chernobyl and Fukushima accidents

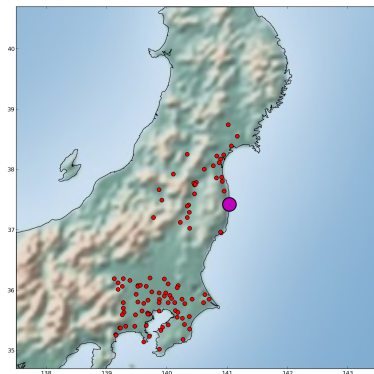


- ▶ 30 deaths in the first days of the accident
- ▶ 200 000 evacuees
- ▶ 30 km exclusion zone
- ▶ Mid and long term sanitary impact: thyroid cancer (thousands of cases).

Case study: Chernobyl and Fukushima accidents

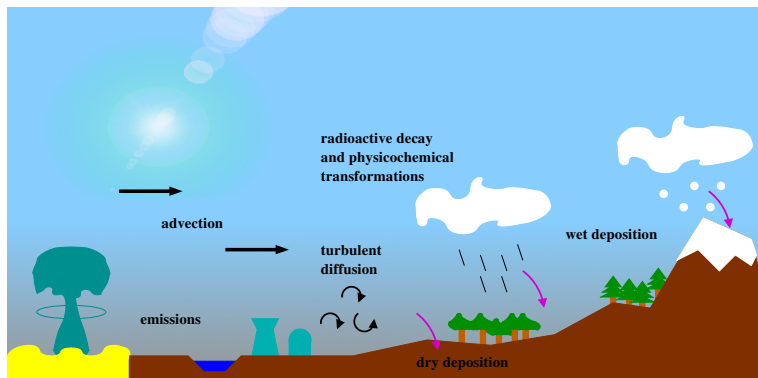


Caesium-137 deposition [IRSN database]



Air quality monitoring network

Case study: Chernobyl and Fukushima accidents



- ▶ Modelled by PDEs of the transport processes and physical and chemical parametrisations.
- ▶ Source term usually unknown.
- ▶ Parameters of the physical parametrisations often poorly known (effective turbulent diffusion, scavenging and dry deposition parameters).

Case study: Cost function

- ▶ Source-receptor relationship: **H**. Linear model.
- ▶ Problem usually solved using 4D-Var [Bocquet, 2012] or methods equivalent to the **representer technique**. Here, study focused on UQ of the *best* estimate [Liu et al., 2017].
- ▶ Log-normal errors for the prior and for the observations. Non-Gaussian statistics.
- ▶ Cost function from Bayes' rule:

$$\begin{aligned} \mathcal{L}(\mathbf{z}; \theta) &= -\ln p(\mathbf{z}|\mathbf{y}, \theta) = -\ln p(\mathbf{y}|\mathbf{z}, \theta) - \ln p(\mathbf{z}|\theta) + \ln p(\mathbf{y}|\theta) \\ &= \frac{1}{2} \|\ln \mathbf{y} - \ln \mathbf{H}\bar{\mathbf{x}}e^{\mathbf{z}}\|_{\mathbf{R}^{-1}}^2 + \frac{1}{2} \|\mathbf{z}\|_{\mathbf{B}^{-1}}^2 + \frac{1}{2} \ln |\mathbf{R}| + \frac{1}{2} \ln |\mathbf{B}| + \xi. \end{aligned}$$

- ▶ Two strategies to quantify the uncertainty of the best estimate:
 - Bayesian hierarchy (HB):

$$p(\mathbf{x}, \theta | \mathbf{y}) = \frac{p(\mathbf{y} | \mathbf{x}, \theta) p(\mathbf{x} | \theta) p(\theta)}{p(\mathbf{y})}, \quad p(\mathbf{x} | \mathbf{y}) = \int d\theta p(\mathbf{x}, \theta | \mathbf{y}). \quad (2)$$

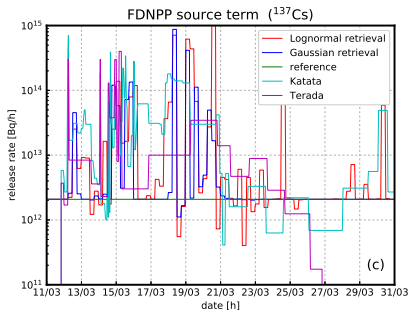
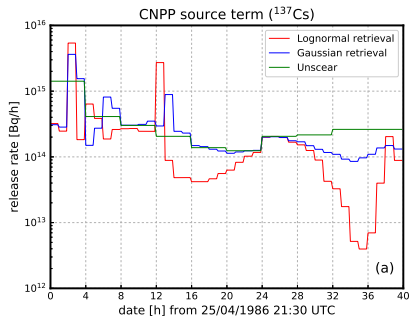
- Empirical Bayes (EB):

$$p(\mathbf{x} | \mathbf{y}) \approx p(\mathbf{x} | \mathbf{y}, \theta^*). \quad (3)$$

θ^* here estimated by the Expectation-Maximisation (EM) algorithm.

Case study: Inversions (EB)

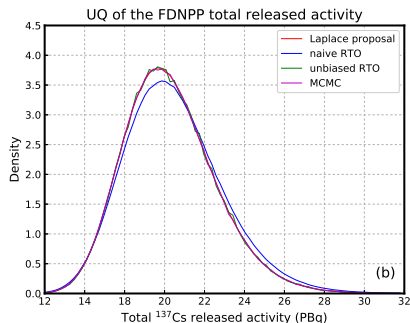
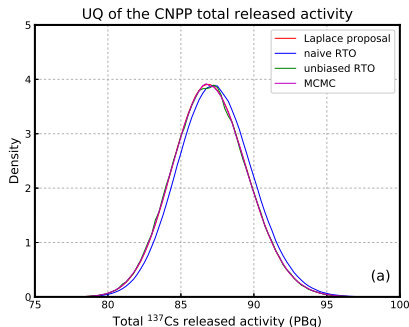
- Uniform hyperparameters: $\mathbf{R} = r^* \mathbf{I}$, $\mathbf{B} = b^* \mathbf{I}$, where r^* and b^* are obtained from EM.



- Chernobyl and Fukushima-Daiichi source terms with Gaussian and lognormal assumptions on the observation errors. Comparison with the Unsear reference source term.

Case study: UQ of the retrieved total radioactivity (EB)

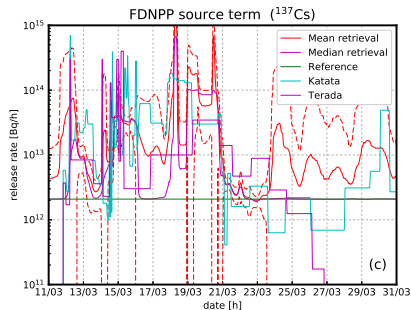
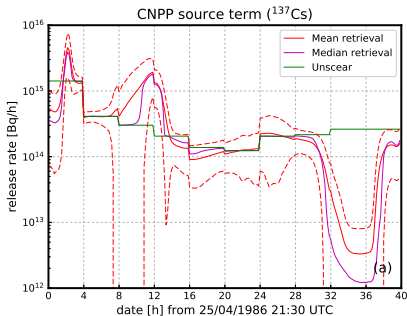
- Probability density function of the total released activity for Chernobyl and Fukushima-Daiichi.



- EB: optimal hyperparameters are first determined. Followed by nonlinear sampling of the total activity using three methods: with a Laplace proposal, a random-then-optimize sampling, an unbiased random-then-optimize sampling and a basic MCMC.

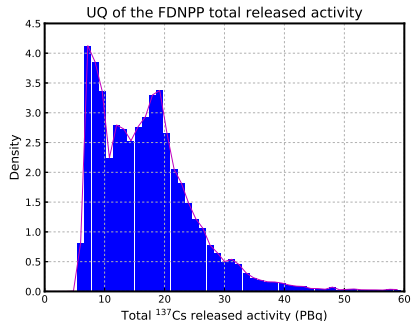
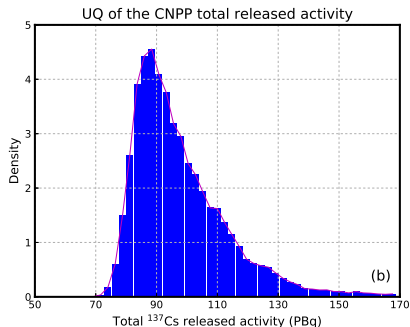
Case study: Inversion (HB)

- ▶ Full solution of the Bayesian hierarchy (HB)
- ▶ Obtained from a Monte Carlo Markov Chain (MCMC)
- ▶ Transdimensional analysis (adaptive grid). Here using only 20 grid cells for Chernobyl and 40 grid cells for Fukushima.



Case study: UQ of the retrieved total radioactivity (HB)

- Probability density function of the total released activity for Chernobyl and Fukushima-Daiichi.



- Full solution of the Bayesian hierarchy; obtained from an MCMC.
- Transdimensional analysis (adaptive grid). Here using only 20 grid cells for Chernobyl and 40 grid cells for Fukushima.

Final word

Thank you for your attention!

- ▶ Part I: A gentle introduction to DA.
- ▶ Part II: More advanced topics including EnKF and EnVar.
- ▶ Part III: Applications of DA including emerging ones such as: glaciology, biology, geomagnetism, medicine, imaging and acoustics, economics and finance, traffic control, etc.

